



Geo-spatial Data in Natural Resources

Homepage: <http://www.gathacognition.com/book/gcb1>
<http://dx.doi.org/10.21523/gcb5>

Online Edition

Geo-spatial
Data in
Natural
Resources

Technical Paper

Chapter 2

Morphometric Analysis of Adula River Basin in Maharashtra, India using GIS and Remote Sensing techniques

Amit Bera^{1*}, Bhabani P. Mukhopadhyay¹, Debasish Das²

1. Department of Earth Sciences, Indian Institute of Engineering Science and Technology, Shibpur, West Bengal, India.
2. Department of Environmental Science, University of Kalyani, West Bengal, India.

Abstract

The present paper attempts to study the morphometric characteristics of Adula River Basin in Maharashtra using GIS and Remote Sensing technique. Topographical maps and ASTER DEM data at 30m spatial resolution have been used for the analysis. Stream networks, different linear, areal, and relief aspects of the basin were analyzed using Arc GIS 10.1 software. The drainage pattern of Adula River is dendritic to sub-dendritic type, with 4th order stream. The mean bifurcation ratio is 5.36, which indicate uniform lithology in the basin. The elongation ratio and circulatory ratio shows elongated shape of the basin. Low length of overland flow (0.63) of the basin is the indicator of moderately low surface runoff. Estimated drainage density (0.79 km/km²) and drainage texture (0.84) of the River basin indicates highly resistant or permeable subsoil material with coarse to moderately coarse drainage texture. The relief aspect of the Adula basin reveals that the major part of the basin has low to moderate relief and values of dissection index and ruggedness index both were very low.

Published Online

27th July 2018

Keywords

Morphometric Analysis;
Remote Sensing;
GIS;
Stream Network;
Adula River.

Editor(s)

R. B. Zolekar

* Author's address for correspondence

Department of Earth Sciences, Indian Institute of Engineering Science and Technology, Shibpur – 711103, West Bengal, India. Mob.: +91 9735774298

Email: amit.rs2017@geology.iiests.ac.in (A. Bera – Corresponding author); bpmbses@gmail.com (B. Mukhopadhyay); ddas_kly@rediffmail.com (D. Das)

<http://dx.doi.org/10.21523/gcb5.1702>

© 2018 GATHA COGNITION® All rights reserved.

1 INTRODUCTION

‘Morphometry may be defined as the measurement and mathematical analysis of the configuration of the earth’s surface and of the shape and dimensions of its landforms’ (Clarke, 1966). Drainage basins are the fundamental units of the fluvial landscape and a great amount of research has focused on their geometric characteristics, which include the topology of the stream network and quantitative description of drainage texture, pattern, shape, and relief characteristics (Abrahams, 1984; Huggett and Cheesman, 2002). A quantitative morphometric characterization of a drainage basin is considered to be the most satisfactory method for the proper planning of watershed management because it enables us to understand the relationship among different aspects of the drainage pattern of the basin, and also to make a comparative evaluation of different drainage basins developed in various geologic and climatic regimes (Zende and Nagrajan, 2011). Fluvial morphometric study of a drainage basin includes the consideration of linear, areal and relief aspects, where the linear aspect deals with the hierarchical orders of streams, numbers and length of stream segments, etc. The areal aspect includes the analysis of basin parameters, basin shape both geometrical and topological (Stream frequency, Drainage density) and the relief aspect includes, the study of absolute and relative relief ratios, average slope, dissection index (Singh, 1998; Khakhlari and Nandy, 2016).

Morphometric parameters mainly depend upon lithology, bed rock and geological structures. Hence, the information of geomorphology, hydrology, geology, and land use pattern is highly informative for reliable study of drainage pattern of the watershed (Astras and Soulankellis, 1992). For quantitative analysis of the watershed involving various components such as stream segments, basin perimeter, basin area, elevation difference, slope and profile of land has been responsible for the natural development of basin (Horton, 1945). Pioneer work on basin morphometry analysis was carried out by Horton (1932, 1945). Then, Horton’s laws were subsequently modified and developed by several geomorphologist and geohydrologist, which was highly appreciable most notably by Strahler (1952), Schumm (1956), Leopold and Miller (1956), Morisawa (1957), Melton (1957), Strahler (1957), Shreve (1967), Muller (1968), Evans (1972), Gregory and Walling (1973), Chorley et al. (1984). In recent decades, the morphometric analysis of the various River basins, have been done by many researchers and scientist (Topaloglu, 2002; Moussa, 2003; Pareta, 2005; Mesa, 2006; Esper, 2008; Magesh et al., 2011; Bhagwat et al., 2011; Wilson et al., 2012; Singh et al., 2014; Sujatha et al., 2014; Meshram and Sharma, 2017; Rai et al., 2017). Gaikwad and Bhagat (2017) have analyzed morphometric parameters for watershed prioritization using multi-criteria analysis with AHP and influence approaches.

Morphometric analysis of River basins using conventional methods is very time consuming, laborious, and cumbersome also. Remote Sensing and GIS is accepted to be powerful geospatial techniques in preparing the drainage map and understanding the watershed's morphometric parameters (Rao *et al.*, 2011). To evaluate the morphometric analysis using Geographical Information Systems (GIS) was attempted by several researchers (Srinivasa *et al.*, 2004; Markose *et al.*, 2014; Das, 2014; Kaliraj *et al.*, 2015; Gajbhiye, 2015; Sahu *et al.*, 2016; Pande and Moharir, 2017; Rai *et al.*, 2017; Chandniha and Kansal, 2017).

Morphometric analysis of streams is an important aspect for characterization of watershed. Proper planning and management of watershed is very necessary for sustainable development (Chandniha and Kansal, 2017). In the present study, an attempt is made to understand the morphometric characteristics of Adula River Basin, a tributary of Pravara River flowing through the Maharashtra using GIS and RS.

2 STUDY AREA

The latitudinal and longitudinal extent of the Adula River basin are between 19°32'40" N to 19°43'2" N and 74°10'15"E to 74°48'18" E. Adula River is one of the major Tributaries of Pravara River (Figure 1). Adula River rises in the north of Akole on the slope of Patta and Mahakali. It flows for fifteen miles in an easterly direction between two ranges of hill which enclose the Samsherpur valley, then falling into the rocky chasm some 150 feet deep. The area as a whole comprises of hill slopes running parallel with the streams on the north and south, pediments extending up to alluvial banks which are deeply dissected to form badlands (Joshi, 2010). The catchment area of Adula River basin is 265.18 square km. Geologically, the area is covered by the thick basaltic lava flows of Deccan Traps. According to Geological Survey of India (GSI, 1976), the region is predominantly covered by pa-hoe-hoe flows. Basaltic rocks and a typical sub rounded weathering products are common in the study area. The soils of this region are covered by thick alluvial soil and black regur soil. The climatic condition of the basin is under the influence of south west monsoon. High amount of precipitation occurs from the month of July to September. The average annual rainfall is 501.8 mm (Joshi, 2010) and maximum humidity was noticed in the month July.

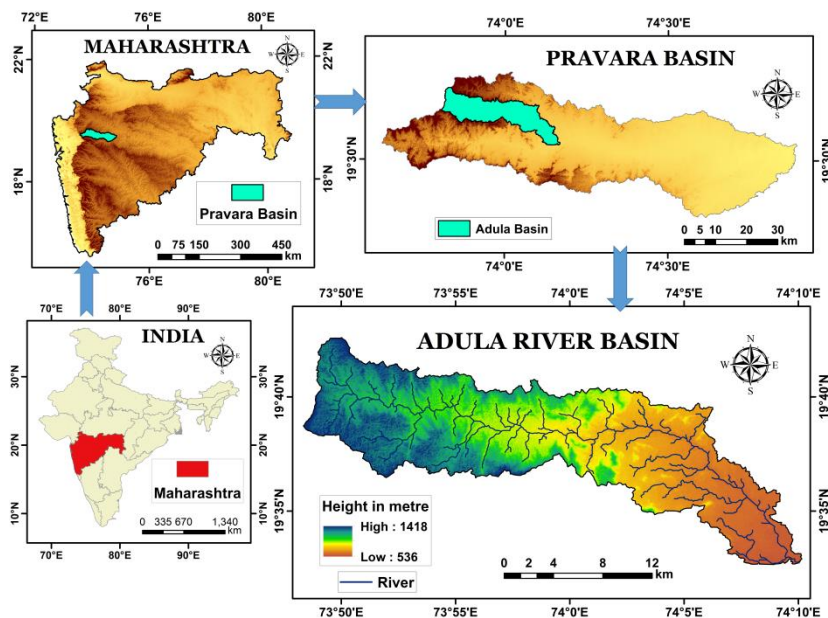


Figure 1. Location map: Adula River basin

3 METHODOLOGY

Survey of India Topo-sheets (47 I/2, 47 E/14) on 1:50,000 and ASTER (Advanced Spaceborne Thermal Emission and Reflection Radiometer) elevation data have been used for generation of Digital Elevation Model. SOI topographic map was georeferenced using WGS 84 datum, Universal Transverse Mercator (UTM) zone 44N projection in ArcGIS desktop 10.1. The Adula watershed was delineated and drainage network was extracted using ASTER DEM (30m resolution) in conjunction with SOI Topo-sheet. Input morphometric maps have been prepared by using GIS Platform Arc GIS 10.1. The methodologies adopted for the computation of selected morphometric parameters are given in [Table 1](#) and [Figure 2](#).

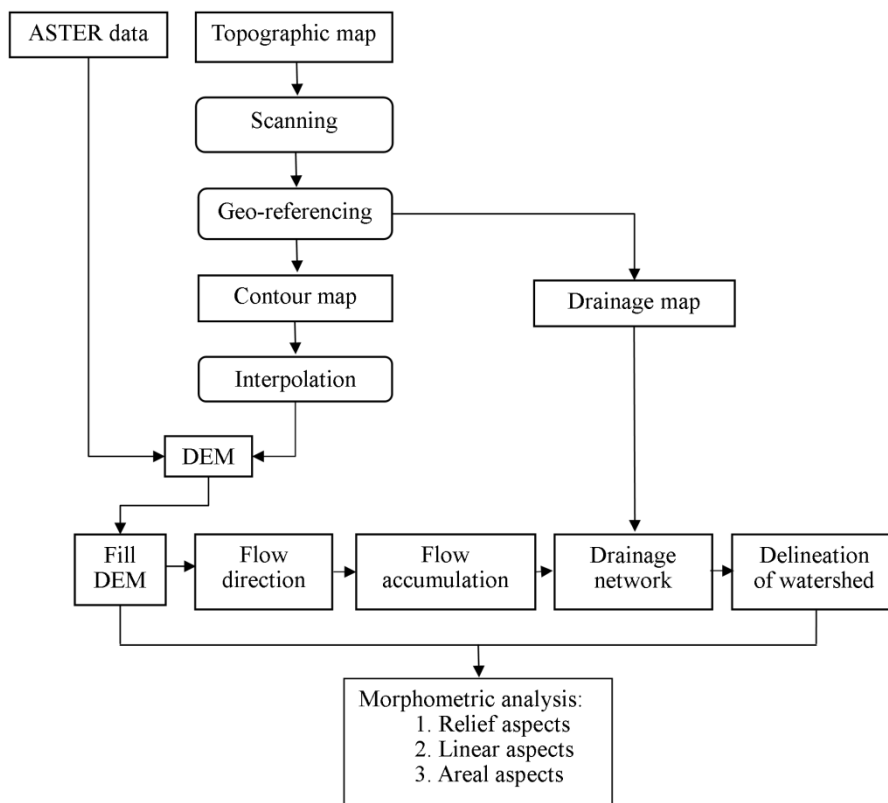


Figure 2. Methodology

4 RESULT AND DISCUSSION

4.1 Linear Aspects

4.1.1 Stream Order (u)

The designation of stream order is the first step in morphometric analysis of a drainage basin, based on hierarchy (Strahler, 1952). The classification of streams based on the number and type of tributary junctions, has proven to be a useful indicator of stream size, discharge, and drainage area (Strahler, 1957). Strahler's method of stream ordering has been adopted and it was found that the Adula River basin is a 4th order drainage basin (Figure 3). The 1st order streams are generally flow during the wet season. Results shows the maximum stream order frequency of the Adula River is observed in case of first-order streams and then for second order.

Table 1. Methodology adopted for computations of morphometric parameters

Parameters	Formula	References
Linear Aspects		
Stream order (u)	Hierarchical rank	Strahler (1952)
Stream length (Lu)	Length of the stream	Horton (1945)
Stream number (Nu)	Nu= Total number of stream of all order	Horton (1945)
Mean stream length (Lsm)	Lsm= Lu/Nu Where, Lu=total stream length of order 'u' Nu= Total number of stream segments of order'u'	Strahler (1964)
Bifurcation ration (Rb)	(Rb) =Nu/Nu+1 Where, Nu= Total stream length of order 'u' and Nu+1=Number of segments of next higher order	Schumm (1956)
Mean bifurcation ratio (Rbm)	Rbm= average of bifurcation ratios of all order	Strahler (1957)
Basin length (Lb)	The straight line from the mouth of the basin to the farthest point on the basin perimeter.	Schumm (1956)
Areal aspect		
Basin area (A)	Area from which water drains to a common stream and boundary determined by opposite ridges.	Strahler (1964)
Basin perimeter (P)	Outer boundary of drainage basin measured in kilometers	Schumm (1956)
Form factor (F _f)	$F_f = A / L_b^2$ Where, A = Basin area, L = Basin length	Horton (1945)
Elongation ratio (Re)	$Re = (2/L_b) * (A / \pi)^{0.5}$ A=Area of the basin,km ² R = Radius of circle whose area equal to basin area , Lb=Basin length	Schumm (1956)

Circularity ratio (Rc)	$Rc = 4 \pi A / P^2$ Where, A= Basin area, $\pi = 3.14$, P = Perimeter of basin	Miller (1953)
Drainage density (Dd)	$Dd = Lu / A$ Where, Lu= Length of the stream , A=Area of basin	Horton (1932)
Stream frequency (Fs)	$Fs = Nu / A$ Where, Nu = Number of stream, A= Basin area	Horton (1932)
Drainage texture (R)	$R = Nu / P$ Where Nu = Number of stream, P= Drainage perimeter	Horton (1945), Smith (1950)
Leminiscate ratio (K)	$K = L^2 / 4A$ Where L = Length of the basin and A = Area of the basin	Schumm (1956)
Constant of channel maintenance (Cc)	$C = 1 / Dd$ Where, Dd= Drainage density	Schumm (1956)
Infiltration number (If)	$If = Fs * Dd$ Where, Fs= Stream frequency Dd= Drainage density	Faniran (1968)
Relief aspect		
Absolute relief	GIS Analysis / DEM	-
Relative relief (Rr)	$Rr = \text{Max.Elevation} - \text{Min. Elevation}$	Smith (1935)
Aspect map	GIS analysis / DEM	-
Dissection index (Di)	$(\text{Max. elevation} - \text{min. Elevation}) / \text{max elevation}$	Nir (1957)
Ruggedness index (Ri)	$Ri = (Rr * Dd) / 1000$ Where, Rr= Relative relief , Dd=Drainage density	Patton and Baker (1976)
Slope map	GIS analysis / DEM	-
Flow direction map	GIS analysis / DEM	-

4.1.2 Stream Number (Nu)

Number of streams is also described as total counts of stream segments of different order separately and is inversely proportional to the stream order (Chandniha and Kansal 2017). Stream numbers of 1st, 2nd, 3rd and 4th order streams are 89, 19, 2 and 1, respectively. The number of streams segments gradually decrease with increase in stream order. High number of 1st order streams some time responsible for sudden flash floods after heavy rainfall.

4.1.3 Stream Length (Lu)

Total length of stream segment in a particular order is represented as stream length (Lu). The each stream length is measured from source to mouth with the help of ArcGIS 10.1, arc tool on the basis of Horton's law of stream length (Horton, 1945) (Table 2).

Table 2. Morphometric analysis: linear aspects

Stream orders (u)	Stream numbers (Nu)	Stream length (Lu) in km	Mean stream length (L _m) in km	Bifurcation ratio (Rb)	Mean bifurcation ratio (R _{b_m})
I	89	115.75	1.30	---	5.36
II	19	43.30	2.28	4.68	
III	2	10.15	5.08	9.5	
IV	1	41.01	41.01	2	

4.1.4 Mean Stream Length (Lsm):

Mean stream length (Lsm) is a characteristic property related to the drainage network components and its associated basin surfaces (Strahler, 1964). It has been calculated by dividing the total stream length of each order by the total number of streams segments in that particular order. In the case of Adula River, the Lsm values exhibit the variation from 1.30 to 41.01.

4.1.5 Bifurcation Ratio (Rb)

The bifurcation ratio is the ratio of the number of stream segments of given order to the number of segments of next higher order (Schumm, 1956). It is a dimensionless property and shows the degree of integration prevailing between streams of various orders in a drainage basin (Magesh et al., 2013). The results showed that the bifurcation ratio is not same for all orders. Geological and lithological development of the drainage basin may be the reason for these variations (Strahler, 1964). Low bifurcation ratio value

indicates poor structural disturbance and the drainage patterns have not been distorted (Strahler, 1964), whereas the high bifurcation ratio value indicates high structural complexity and low permeability of the terrain (Magesh et al., 2013). Bifurcation ratio shows only a small variation for different regions on different environment except where powerful geological control dominates (Strahler, 1957). The value of bifurcation ratio of Adula River varies from 2 to 9.5 (Table 2).

4.1.6 Mean Bifurcation Ratio (R_{bm})

Mean Bifurcation Ratio is the average of bifurcation ratios of all stream orders of a drainage system. Here, the mean bifurcation ratio of Adula River is 5.36 which indicates that there have some strong structural and morphological control over the drainage development (Table 2).

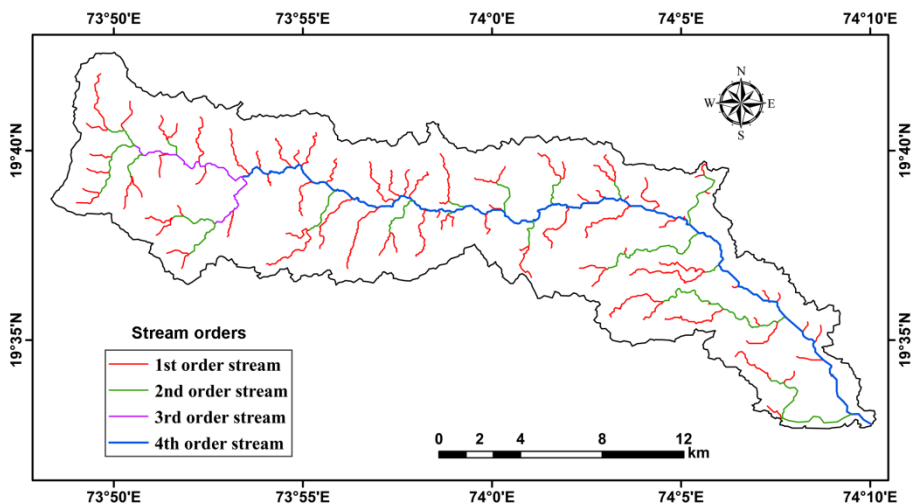


Figure 3. Stream orders

4.2 Areal Aspects

4.2.1 Form Factor (F_f)

Form factor is the numerical index which is commonly used to represent different basin shapes (Horton, 1932). The value of form factor is generally varies between 0.1-0.8. Smaller the value of form factor indicates elongated nature of the basin while higher value is the characteristics of a circular type basin. A near perfect circular basin has a form factor value of near to 0.8. In the case of River Adula the form factor value is 0.16 (Table 3), which indicating elongated shape of the basin.

4.2.2 Elongation Ratio (Re)

Schumm (1956) defined it as the ratio between the diameter of the circle of the same area as the drainage basin and the maximum length of the basin. The value of elongation ratio ranges from 0 to 1. Values close to '1' have very low relief with circular shape while values close to '0' associated with high relief and steep slope with elongated shape. The elongation ratio of Adula River is 0.46 which indicates basin to be moderately elongated with moderately high relief and steep slope.

Table 3. Morphometric analysis: areal aspects

Morphometric Parameters	Result
Basin area	265.18 (km ²)
Basin perimeter	131.61 (km)
Form factor (F_f)	0.16
Elongation ratio (Re)	0.46
Circularity ratio (R_c)	0.19
Drainage density (D_d)	0.79 (km/km ²)
Stream frequency (F_s)	0.42/ km ²
Leminiscate ratio (K)	0.08
Drainage texture (R)	0.84
Constant of channel maintenance (C_c)	1.26 (km ² /km)
Infiltration number (I_f)	0.33
Length of overland flow (L_g)	0.63

4.2.3 Circularity Ratio (R_c)

It is the ratio of the basin area to the area of a circle having circumference as the perimeter of the basin. Circularity ratio is dimensionless and expresses the degree of circularity of the basin (Miller, 1953). Its value ranges between 0 to 1 and values close to 1.0 represent the more or less circular shape of the basin with structurally controlled. The circulatory ratio is mainly concerned with the length and frequency of streams, geological structures, land use land cover, climate, relief and slope of the basin. It is a significant ratio that indicates the dendritic stage of a watershed. Low, medium and high values of circularity ratio show the young, mature, and old phases of the life cycle of the tributary watershed (Wilson et al., 2012; Rai et al., 2017). The value of circularity ratio of Adula River basin is 0.19, which is the indication of elongated shape of basin with highly permeable homogenous geologic materials.

4.2.4 Drainage Density (Dd)

Drainage density is one of the important indicators of the landform element and provides a numerical measurement of landscape dissection and runoff potential (Chorley, 1969). It is defined as the total stream length in a given basin to the total area of the basin (Strahler, 1964). The drainage density in the study area is critically low (Figure 4). The drainage density of the study area has been classified into nine different textures. Comparatively high drainage density ($>1.28 \text{ km/km}^2$) was found at the central and the eastern most part of the basin. The main localities under this category are Samshepur, Tahakari, Hivargaon etc. Low and moderate texture drainage density was found at the both northern and southern most part of the basin.

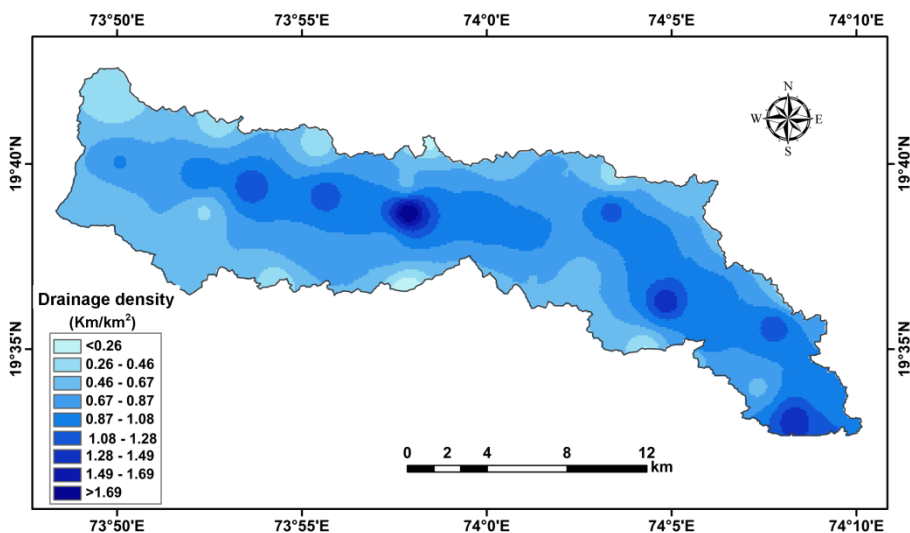


Figure 4. Drainage density

4.2.5 Stream Frequency (Fs)

The total number of stream segments of all orders per unit area is known as stream frequency (Horton, 1932). Stream frequency shows a positive association with drainage density suggesting an increase in stream population with increasing drainage density. The total number of stream segments of all orders per unit area is known as stream frequency (Horton, 1932). Stream frequency shows a positive association with drainage density suggesting an increase in stream population with increasing drainage density. Stream frequency mainly depends on the lithology of the basin and reflects the texture of the drainage network, infiltration capacity, vegetation cover, relief, nature and amount of rainfall and subsurface material permeability (Parveen et al.,

2012; Damilola, 2016). The stream frequency has been classified into eight different classes (Figure 5). The mean stream frequency for the whole Adula River basin is $0.42/\text{km}^2$, which indicates low stream frequency. The low stream frequency indicates low degree of dissection and vice versa.

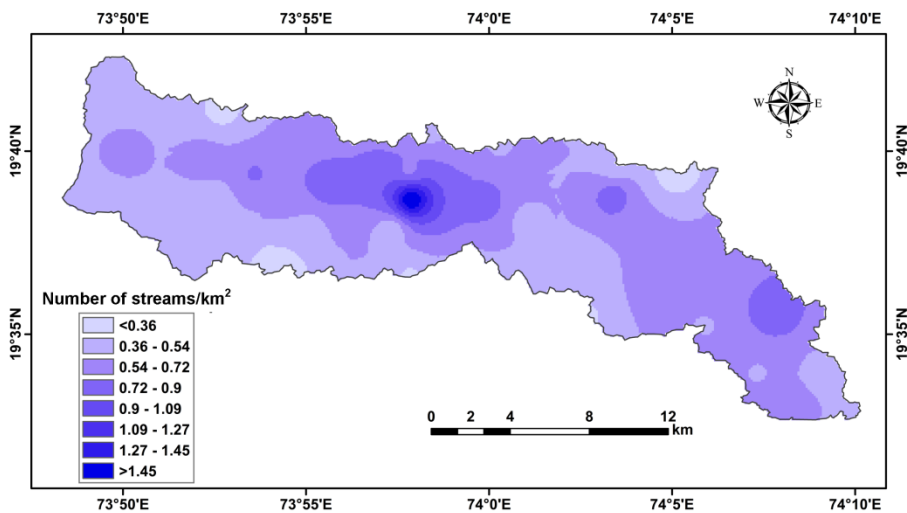


Figure 5. Stream frequency

4.2.6 Leminiscate Ratio (*K*)

Leminiscate ratio expresses the sloping characteristics of a River basin. Generally Lower value of Leminiscate ratio indicates lower intensity of runoff and vice-versa. Low values represent basin nearly rounded and prevailing vertical and lateral erosions, which refer to geomorphic stage of development for a basin (Ashour and Torab, 1991). The calculated leminiscate ratio of the Adula River basin is 0.08, which indicate the runoff was critically low in the basin.

4.2.7 Drainage Texture (*R*)

Drainage Texture (*R*) is one of the significant concept in River morphometry which means that the relative spacing of stream lines. Drainage texture is the product of number of stream and drainage perimeter. Low drainage density is the indication of coarse drainage texture while high drainage density indicates fine drainage texture. Drainage texture of Adula River basin is 0.84, which is indication of coarse texture river basin.

4.2.8 Constant of Channel Maintenance (C_c)

The constant of channel maintenance indicates the relative size of landform units in a drainage basin and has a specific genetic connotation (Strahler, 1957). It is depending on different types of geomorphological characteristics like, rock type permeability, climatic regime, vegetation etc. Constant of channel maintenance of Adula river basin is $1.26 \text{ km}^2/\text{km}$.

4.2.9 Infiltration Number (I_f)

Infiltration number (I_f) of a River basin is the product of drainage density and stream frequency, which helps to understand the infiltration characteristics of the basin. It provides a significant idea about the infiltration characteristics of River basin area. It is inversely proportional to the infiltration capacity of the basin. The higher the infiltration number, the lower will be the infiltration and the higher run-off (Rai et al., 2017). The infiltration number of Adula River basin is 0.33, which indicating that the infiltration capacity is very high resulting in very low runoff.

4.2.10 Length of Overland Flow (L_g)

Length of overland flow (L_g) is one of the most vital self-governing variables affecting hydrologic and physiographic development of drainage basin (Horton, 1932; Javed et al., 2011). It is inversely proportional to half of reciprocal of drainage density and is the length of water over the ground before it gets concentrated into definite stream channels (Horton, 1945). The length of overland flow of the Adula River basin is 0.63 (Table 3), which shows moderately low surface runoff of the basin.

4.3 Relief Aspects

4.3.1 Absolute Relief

Absolute relief contains maximum altitude at each point. Absolute relief of entire Adula River basin is divided into five categories (Figure 6). High absolute relief ($>997\text{m}$) was found in the western most part of the basin, basically in the upper catchment area. The absolute relief is gradually decreases towards the mouth of the river. Central portion of the basin and some small pockets shows the moderate type of absolute relative relief.

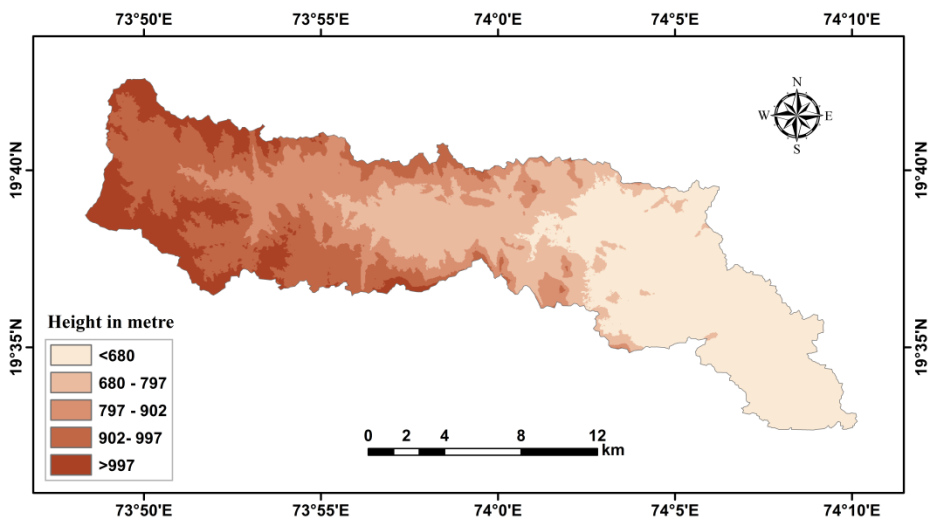


Figure 6. Absolute relief

4.3.2 Aspect Map

Aspect map generally refers to the direction to which a mountain slope faces. The aspect map is a very important parameter to understand impact of sun on local climate of the area. Generally west facing slope showing the hottest time of day in the afternoon and in most cases a west-facing slope will be warmer than sheltered an east-facing slope (Singh et al., 2014). Aspect has significant effects on growth and type's vegetation distribution of an area. The aspect map of the Adula River Basin was derived from ASTER DEM. The slope aspects of entire basin is divided into ten categories: Flat-1, North- 0-22.50, Northeast- 22.5-67.50, East- 67.5-112.50, Southeast- 112.5-157.50, South- 157.5-202.50, Southwest- 202.5-247.50, West- 247.5-292.50, Northwest- 292.5-337.50, North- 337.5-360. Principal stream of the area is flowing from north-west to south-east direction. Aspect map (Figure 7) of Adula River basin indicates that the west to south-west facing slope are dominant along the left bank of the principal.

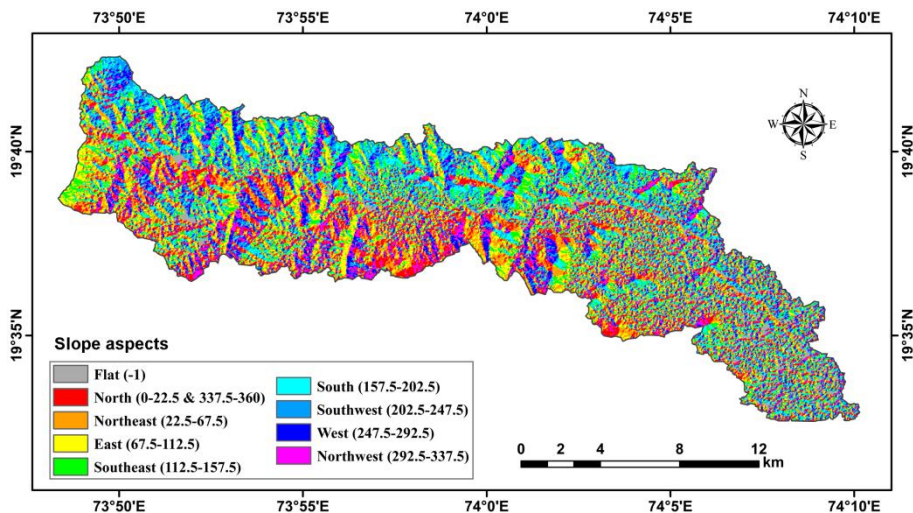


Figure 7. Slope aspects

4.3.3 Relative Relief (R_r)

The relative relief (R_r) represents actual variation of altitude in a unit area with respect to its local base level. The relative relief does not take into account the dynamic potential of the terrain but as it is closely associated with slopes and it is more expressive and also useful in understanding the morphogenesis of this region (Bhunia et al., 2012). Relative relief of entire study area is divided into five categories (Figure 8). High relative relief (>205m) was found in the upper most part of the basin and some small pockets of high relative relief were also illustrated in the western and central part of the basin. Moderately relative relief (158 -205m) found in the central and western most part of the basin. Lowest relative relief (<105m) was found in the eastern most part of the basin.

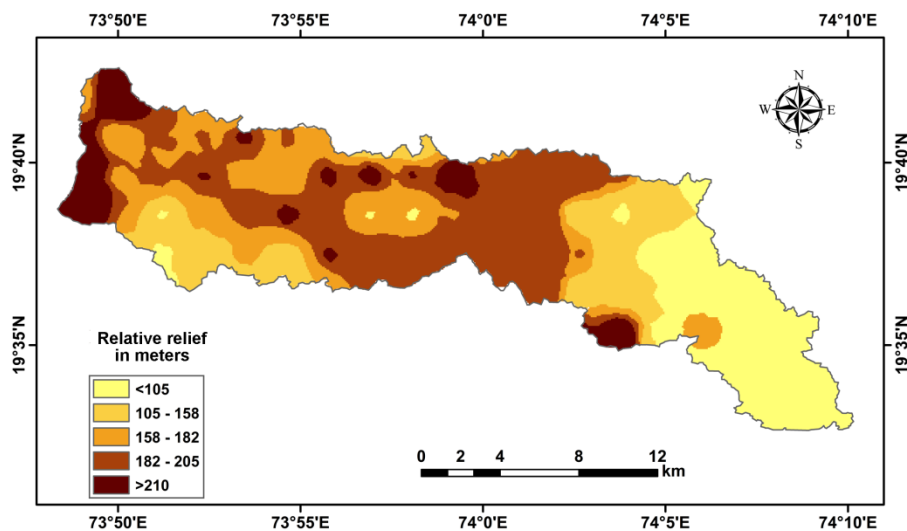


Figure 8. Relative relief

4.3.4 Slope

The slope angle indicates the magnitude of the component of the gravitational surface acting to produce movement of solid bodies, water or soil particles down a slope (Strahler, 1956). Slope is a significant morphological parameter which directly controls infiltration and runoff of any terrain. The slope map for Adula River basin (Figure 9) was developed with the help of DEM using Arc GIS Spatial Analyst tools. About 83.83% area of Adula River basin comes under gentle and moderate slope (<14.54), which indicates nearly flat terrain and most favorable for infiltration.

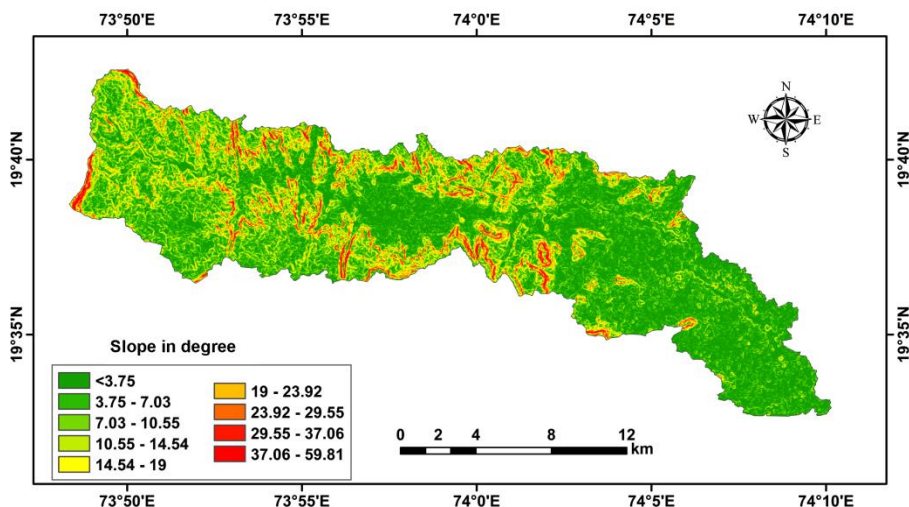


Figure 9. Slope

4.3.5 Dissection Index (Di)

Dissection index (Di) is expressing the ratio of the maximum relative relief to maximum absolute relief. It is an important morphometric indicator of the nature and magnitude of dissection of terrain (Singh, 2000). The value of dissection index varies from zero to one, where zero indicates complete absence of dissection and one indicates vertical cliff. Generally, low 'Di' corresponds with the subdued relief or old stage, and with low relative relief. Conversely, the areas with high 'Di' indicate high relative relief where slope of the land is steep (Deen, 1982). The dissection index, for the study area varies from 0.18 to 0.49. It has been classified into five categories (Figure 10). Very low (<0.18) dissection index has been found mostly in the eastern part of the basin i.e. Chikhali, Rajapur, Biroba wasti, Dongargaon area. Moderate to moderately high (0.18-0.28) dissection index was found Northwestern part of the basin. High dissection index (>0.44) found at the central and the western most part of the basin. The main localities under this category are Deothan, Kokanwadi, Chandgirwadi, etc.

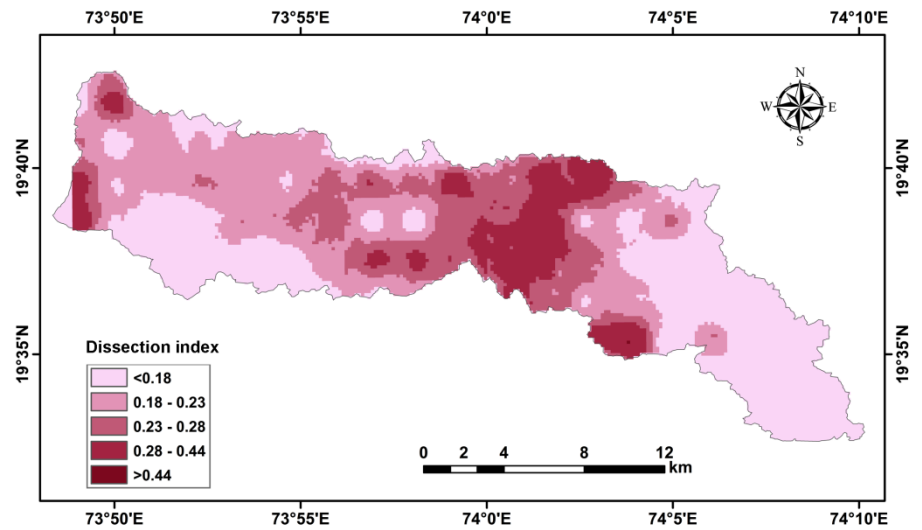


Figure 10: Dissection index

4.3.6 Ruggedness Index

Ruggedness Index indicates roughness of the terrain and the degree of dissection of topography where drainage has considered as an important parameter. This index is being widely used by the earth scientists in connection with the morphological studies of terrain and it leads to better understanding of the surface configuration evolved under complex geomorphic processes. This index reflects the combined effects of evolutionary rhythmic processes in the development of relief (Prakasam and Biswas, 2012; Mukhopadhyay, 1984). Ruggedness index (Figure 11) varies from 0.04 to 0.21.

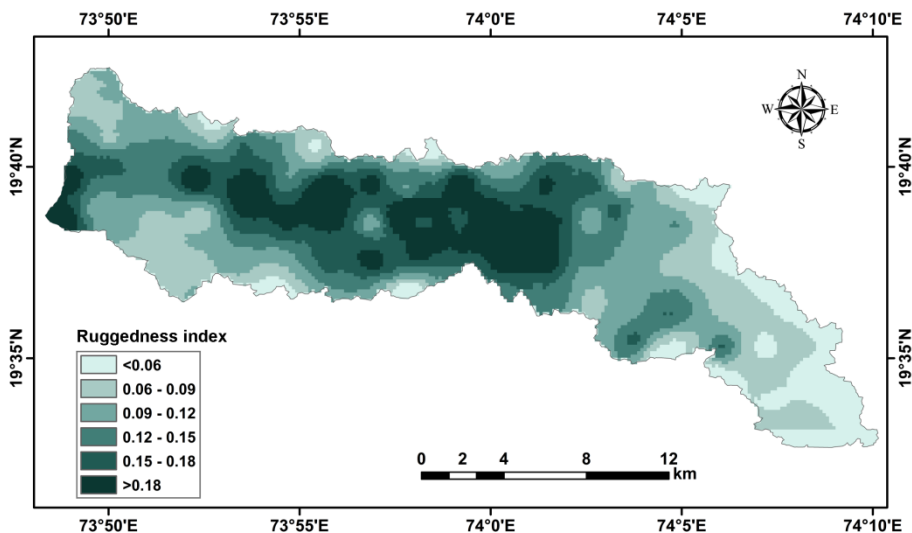


Figure 11. Ruggedness index

5 CONCLUSION

Morphometric analysis of drainage basins thus provides not only an elegant description of the landscape, but also serve as a powerful means of comparing the form and process of drainage basins that may be widely separated in space and time (Easterbrook, 1993). The linear and areal aspect of Adula River basin reveals that the basin is elongated shape with highly permeable homogenous geologic materials. From the relief parameter of the Adula River basin, it is inferred that major part of the basin has low to moderately high relief. The results of morphometric parameters of Adula River Basin can be used for integrated watershed management, groundwater recharge zone mapping and rainwater harvesting.

CONFLICT OF INTEREST

The authors declare no conflict of interest.

ACKNOWLEDGEMENT

The authors thankful to the [Indian Institute of Remote Sensing \(IIRS\)](#), Indian Space Research Organization (ISRO) for continue support during the work. The authors also extend their thanks to anonymous reviewers for the valuable comments and suggestions.

REFERENCES

Abrahams, A. D., 1984. Channel Networks: A Geomorphological Perspective. *Water Resources Research*, 20(2), 161-188. DOI: 10.1029/wr020i002p00161

- Ashour, M. M., and Torab, M. M., 1991. Morphometric Analysis of Basins and Drainage Networks: In Goda, H. *Handbook of Morphometric Analysis*, 267-376.
- Astras, T., and Soulanellis, N., 1992. Contribution of digital image analysis techniques on Landsat-5 TM imageries for drainage delineation. A case study from the Olympus mountain, west Macedonia, Greece. In *Proc. 18th Annual Conf. Remote Sensing Soc., Univ of Dundee*, 163-172.
- Bhagwat, T. N., Shetty, A., and Hegde, V. S., 2011. Spatial variation in drainage characteristics and geomorphic instantaneous unit hydrograph (GIUH); implications for watershed management-A case study of the Varada River basin, Northern Karnataka. *Catena*, 87(1), 52-59. DOI: 10.1016/j.catena.2011.05.007
- Bhunia, G. S., Samanta, S. and Pal, B., 2012. Quantitative analysis of relief characteristics using space technology. *International Journal of Physical and Social Sciences*, 2(8), 350-365.
- Chandniha, S. K., and Kansal, M. L., 2017. Prioritization of sub-watersheds based on morphometric analysis using geospatial technique in Piperiya watershed, India. *Applied Water Science*, 7(1), 329-338. DOI: 10.1007/s13201-014-0248-9
- Chorley, R. J., 1969. *Introduction to physical hydrology*. Methuen and Co. Ltd., Suffolk, 211.
- Chorley, R. J., Schumm, S. A. and Sugden, D. E., 1984. *Geomorphology: London*. Methuen and Co. London.
- Clarke, J. I., 1966. *Morphometry from Maps, Essays in geomorphology*. Elsevier Publ. Co., New York, 235-274.
- Damilola, A. E., 2016. The Osun Drainage Basin in the Western Lithoral Hydrological Zone of Nigeria: A morphometric study. *GEOGRAFIA, Malaysian Journal of Society and Space*, 12(8), 71-88.
- Das, D., 2014. Identification of Erosion Prone Areas by Morphometric Analysis Using GIS. *Journal of the Institution of Engineers (India): Series A*, 95(1), 61-74. DOI:10.1007/s40030-014-0069-8
- Deen, M., 1982. *Geomorphology and Land use: a Case Study of Mewat*. Doctoral Dissertation, Jawaharlal Nehru University, New Delhi, India.
- Easterbrook, D. J., 1999. *Surface processes and landforms*, Macmillian Publishing Co., New York, 325.
- Esper Angillieri, M. Y., 2008. Morphometric analysis of Colangüil river basin and flash flood hazard, San Juan, Argentina. *Environmental Geology*, 55(1), 107-111. DOI: 10.1007/s00254-007-0969-2
- Evans, I. S., 1972. General geomorphometry, derivatives of altitude, and descriptive statistics. In: Chorley RJ (ed) *Spatial analysis in geomorphology*. Harper and Row, New York, 17-90.
- Gaikwad, R. and Bhagat, V., 2017. Multi-Criteria Watershed Prioritization of Kas Basin in Maharashtra India: AHP and Influence Approaches. *Hydrospatial Analysis*, 1(1), 41-61. DOI: 10.21523/gcj3.17010105
- Gajbhiye, S., 2015. Morphometric Analysis of a Shakkar River Catchment Using RS and GIS. *International Journal of u- and e-Service, Science and Technology*, 8(2), 11-24. DOI: 10.14257/ijunesst.2015.8.2.02
- Gregory, K. J., and Walling, D. E., 1976. *Drainage basin form and process: A geomorphological approach*. Edward Arnold, London.
- GSI, 1976. Geological survey of India: know your District, 125th anniversary publication.
- Horton, R. E., 1945. Erosional development of streams and their drainage basins; hydrophysical approach to quantitative morphology. *Geological society of America bulletin*, 56(3), 275-370.
- Horton, R. E., 1932. Drainage-basin characteristics. *Transactions, American Geophysical Union*, 13(1), 350-361. DOI: 10.1029/tr013i001p00350

- Huggett, R. J. and Cheesman, J., 2002. *Topography and the Environment*. Prentice Hall, Pearson Education, Upper Saddle River, NJ, 274.
- Javed, A., Khanday, M. Y. and Rais, S., 2011. Watershed prioritization using morphometric and land use/land cover parameters: A remote sensing and GIS based approach. *Journal of the Geological Society of India*, 78(1), 63-75. DOI: 10.1007/s12594-011-0068-6
- Joshi, V. U., 2010. Evaluation of geomorphic and anthropogenic effects on the development of badland along adula and mahalingi nadi, pravara basin Maharashtra. A research project supported by UGC.
- Kaliraj, S., Chandrasekar, N. and Magesh, N. S., 2015. Morphometric analysis of the River Thamirabarani sub-basin in Kanyakumari District, South west coast of Tamil Nadu, India, using remote sensing and GIS. *Environmental Earth Sciences*, 73(11), 7375-7401. DOI: 10.1007/s12665-014-3914-1
- Khakhlari, M. and Nandy, A., 2016. Morphometric Analysis of Barapani River Basin in Karbi Anglong District, Assam. *International Journal of Scientific and Research Publications*, 6(10), 238-249.
- Leopold, L. B. and Miller, J. P., 1956. *Ephemeral streams; hydraulic factors and their relation to the drainage net*. US Government Printing Office, USA.
- Magesh, N. S., Chandrasekar, N., and Soundranayagam, J. P., 2011. Morphometric evaluation of Papanasam and Manimuthar watersheds, parts of Western Ghats, Tirunelveli district, Tamil Nadu, India: a GIS approach. *Environmental Earth Sciences*, 64(2), 373-381. DOI: 10.1007/s12665-010-0860-4
- Magesh, N. S., Jitheshlal, K. V., Chandrasekar, N. and Jini, K. V., 2013. Geographical information system-based morphometric analysis of Bharathapuzha river basin, Kerala, India. *Applied Water Science*, 3(2), 467-477. DOI: 10.1007/s13201-013-0095-0
- Markose, V. J., Dinesh, A. C. and Jayappa, K. S., 2014. Quantitative analysis of morphometric parameters of Kali River basin, southern India, using bearing azimuth and drainage (bAd) calculator and GIS. *Environmental earth sciences*, 72(8), 2887-2903. DOI:10.1007/s12665-014-3193-x
- Melton, M. A., 1957. *An analysis of the relations among elements of climate, surface properties, and geomorphology*, CU-TR-11, Columbia Univ New York.
- Mesa, L. M., 2006. Morphometric analysis of a subtropical Andean basin (Tucumán, Argentina). *Environmental Geology*, 50(8), 1235-1242. DOI: 10.1007/s00254-006-0297-y
- Meshram, S. G. and Sharma, S. K., 2017. Prioritization of watershed through morphometric parameters: a PCA-based approach. *Applied Water Science*, 7(3), 1505-1519. DOI:10.1007/s13201-015-0332-9
- Miller, V. C., 1953. *A quantitative geomorphic study of drainage basin characteristics in the Clinch mountain area, Virginia and Tennessees*. Department of Navy, Office of Naval Research. Technical Report 3, Project NR 389-042, Washington DC.
- Morisawa, M., 1957. Accuracy of determination of stream lengths from topographic maps. *Transactions, American Geophysical Union*, 38(1), 86-88. DOI: 10.1029/tr038i001p00086
- Moussa, R., 2003. On morphometric properties of basins, scale effects and hydrological response. *Hydrological Processes*, 17(1), 33-58. DOI: 10.1002/hyp.1114
- Mueller, J. E., 1968. An introduction to the hydraulic and topographic sinuosity indexes. *Annals of the Association of American Geographers*, 58(2), 371-385. DOI: 10.1111/j.1467-8306.1968.tb00650.x
- Mukhopadhyay, S. C., 1984. *The Tista Basin: a study in fluvial geomorphology*. KP Bagchi. Calcutta.

- Pande, C. B. and Moharir, K., 2017. GIS based quantitative morphometric analysis and its consequences: a case study from Shanur River Basin, Maharashtra India. *Applied Water Science*, 7(2), 861-871. DOI: 10.1007/s13201-015-0298-7
- Panhalkar, S. S., 2013. *Rain water harvesting and watershed management in Dudhganga basin Maharashtra*. Doctoral dissertation, Shivaji University, India.
- Pareta, K., 2005. Rainfall-Runoff Modelling, Soil Erosion Modelling, Water Balance Calculation, and Morphometric Analysis of Molali Watershed, Sagar, Madhya Pradesh using GIS and Remote Sensing Techniques. In *25th International Cartographic Congress, INCA*.
- Parveen, R., Kumar, U. and Singh, V. K., 2012. Geomorphometric Characterization of Upper South Koel Basin, Jharkhand: A Remote Sensing & GIS Approach. *Journal of Water Resource and Protection*, 04(12), 1042-1050. DOI: 10.4236/jwarp.2012.412120
- Prakasam, C. and Biswas, B., 2012. Evaluation of geomorphic resources using GIS technology: a case study of selected villages in Ausgram Block, Burdwam District, West Bengal, India. *International Journal of Geology Earth and Environmental Sciences*, 2(2), 193-205.
- Rai, P. K., Mishra, V. N. and Mohan, K., 2017. A study of morphometric evaluation of the Son basin, India using geospatial approach. *Remote Sensing Applications: Society and Environment*, 7, 9-20. DOI: 10.1016/j.rsase.2017.05.001
- Rai, P. K., Chaubey, P. K., Mohan, K. and Singh, P., 2017. Geoinformatics for assessing the inferences of quantitative drainage morphometry of the Narmada Basin in India. *Applied Geomatics*. 1-23. DOI: 10.1007/s12518-017-0191-1
- Rai, P. K., Mohan, K., Mishra, S., Ahmad, A. and Mishra, V. N., 2017. A GIS-based approach in drainage morphometric analysis of Kanhar River Basin, India. *Applied Water Science*, 7(1), 217-232. DOI: 10.1007/s13201-014-0238-y
- Rao, L. A., Rehman, A. Z. and Alia, Y., 2011. Morphometric analysis of drainage basin using remote sensing and GIS techniques: a case study of Etmadpur Tehsil, Agra District. *International Journal of Research in Chemistry and Environment*, 1(2), 36-45.
- Sahu, N., Obi Reddy, G. P., Kumar, N., Nagaraju, M. S., Srivastava, R. and Singh, S. K., 2016. Morphometric analysis in basaltic Terrain of Central India using GIS techniques: a case study. *Applied Water Science*, 1-7. DOI: 10.1007/s13201-016-0442-z
- Schumm, S. A., 1956. Evolution of drainage systems and slopes in badlands at Perth Amboy, New Jersey. *Geological Society of America Bulletin*, 67(5), 597-646. DOI: 10.1130/0016-7606(1956)67[597:eodsas]2.0.co;2
- Shreve, R. L., 1967. Infinite Topologically Random Channel Networks. *The Journal of Geology*, 75(2), 178-186. DOI:10.1086/627245
- Singh, P., Gupta, A., and Singh, M., 2014. Hydrological inferences from watershed analysis for water resource management using remote sensing and GIS techniques. *The Egyptian Journal of Remote Sensing and Space Science*, 17(2), 111-121. DOI: 10.1016/j.ejrs.2014.09.003
- Singh, S., 1998. *Geomorphology*. Allahabad: Prayag Pustak Bhavan, 353.
- Singh, S., 2000. *Geomorphology*, Prayag Pustak Bhavan, Allahabad, 642.
- Srinivasa, V., S., Govindaiah, S. and Honne Gowda, H., 2004. Morphometric analysis of sub-watersheds in the Pavagada area of Tumkur district, South India using remote sensing and GIS techniques. *Journal of the Indian Society of Remote Sensing*, 32(4), 351-362. DOI: 10.1007/bf03030860
- Strahler, A. N., 1952. Hypsometric (area-altitude) analysis of erosional topography. *Geological Society of America Bulletin*, 63(11), 1117-1142. DOI: 10.1130/0016-7606(1952)63[1117:haoet]2.0.co;2
- Strahler, A. N., 1956. Quantitative slope analysis. *Geological Society of America Bulletin*, 67(5), 571-596. DOI: 10.1130/0016-7606(1956)67[571:qsa]2.0.co;2

- Strahler, A. N., 1964. Quantitative geomorphology of drainage basins and Channel networks. In: *Hand book of Applied Hydrology*, Chow, V. T. (Ed.), 4, 39- 44.
- Strahler, A. N., 1957. Quantitative analysis of watershed geomorphology. *Transactions, American Geophysical Union*, 38(6), 913-920. DOI: 10.1029/tr038i006p00913
- Sujatha, E. R., Selvakumar, R. and Rajasimman, B., 2014. Watershed prioritization of Palar sub-watershed based on the morphometric and land use analysis. *Journal of Mountain Science*, 11(4), 906-916. DOI: 10.1007/s11629-012-2628-7
- Topaloğlu, F., 2002. Estimation of instantaneous peak flows in Seyhan river basin using regional regression procedures. *Turkish Journal of Agriculture and Forestry*, 26(1), 47-55.
- Wilson, J. S., Chandrasekar, N. and Magesh, N.S., 2012. Morphometric analysis of major sub-watersheds in Aiyar and Karai Pottanar Basin, central Tamil Nadu, India using remote sensing and GIS techniques. *Bonfring International Journal of Industrial Engineering and Management Science*, 2 (1), 8-15.
- Zende, A. N. and Nagrajan, R., 2011. Drainage Morphology Approach for Water Resources Development of Sub-Watershed in Krishna Basin. *International Journal of Computer & Communication Technology*, 2(8), 13-21.
